



15 - 19 June 2026
Kraków, Poland

FRINGE 2026



Kinematics, Dynamics, and Seismic Potential of the Tianshan from Sentinel-1 InSAR & GNSS

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18 June 2026



Outline

1

Interseismic Velocity and Strain Rate

2

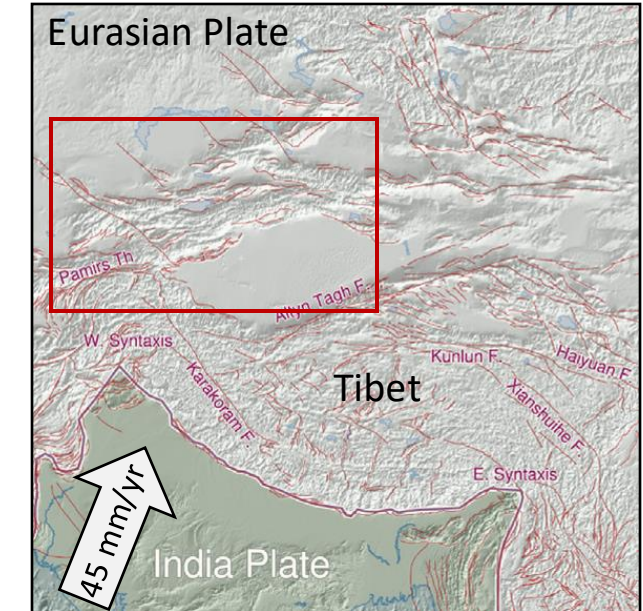
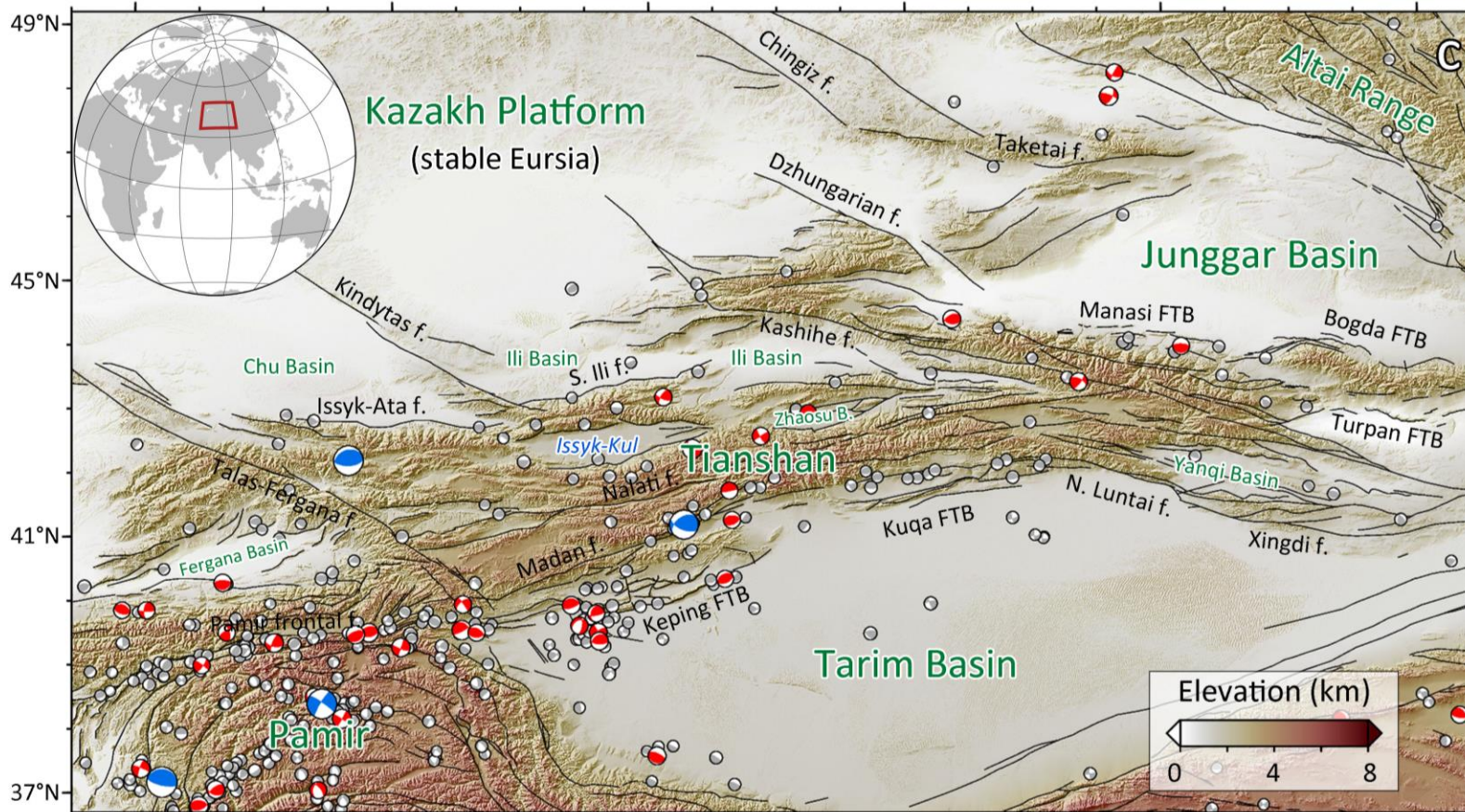
Slip Deficit Rate and Seismic Potential

3

From Kinematics to Dynamics



Tectonic Background



Motivation - Science & Society

- Tectonically active (~ 20 mm/yr, about half the plate convergence rate)
- Frequent earthquakes (over 100 $M_w \geq 6$ earthquakes)
- Kinematics, Dynamics, and Seismic Risk
- Resources and Energy (Oil/Gas/Minerals)
- Infrastructure (Energy pipelines/Railway)

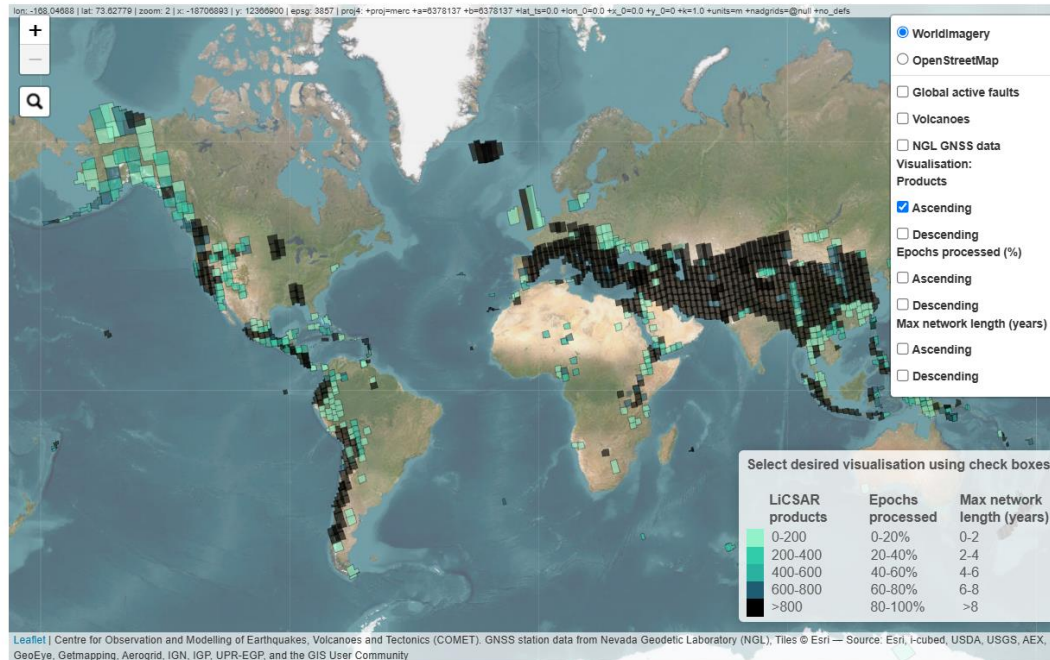
InSAR Time Series Analysis



COMET-LiCS Sentinel-1 InSAR portal



Last compiled on 2026-06-15. Total interferograms: 2,511,362. Change since last month: 17,403

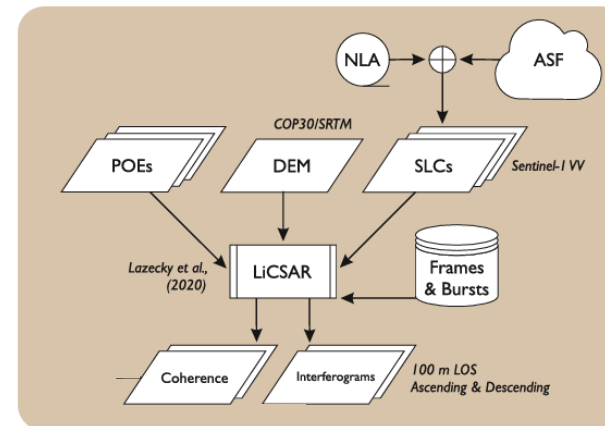


LiCSAR data: 21 tracks/92 frames/~100,000 ifgs/2014-2024

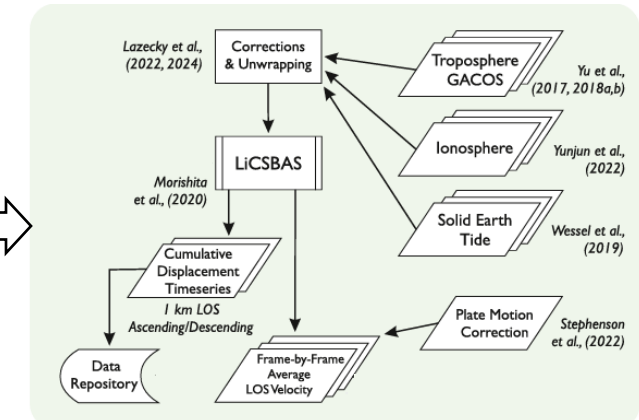
(Lazecký et al., 2020)

Time Series Analysis - LiCSBAS (Morishita et al., 2020)

InSAR



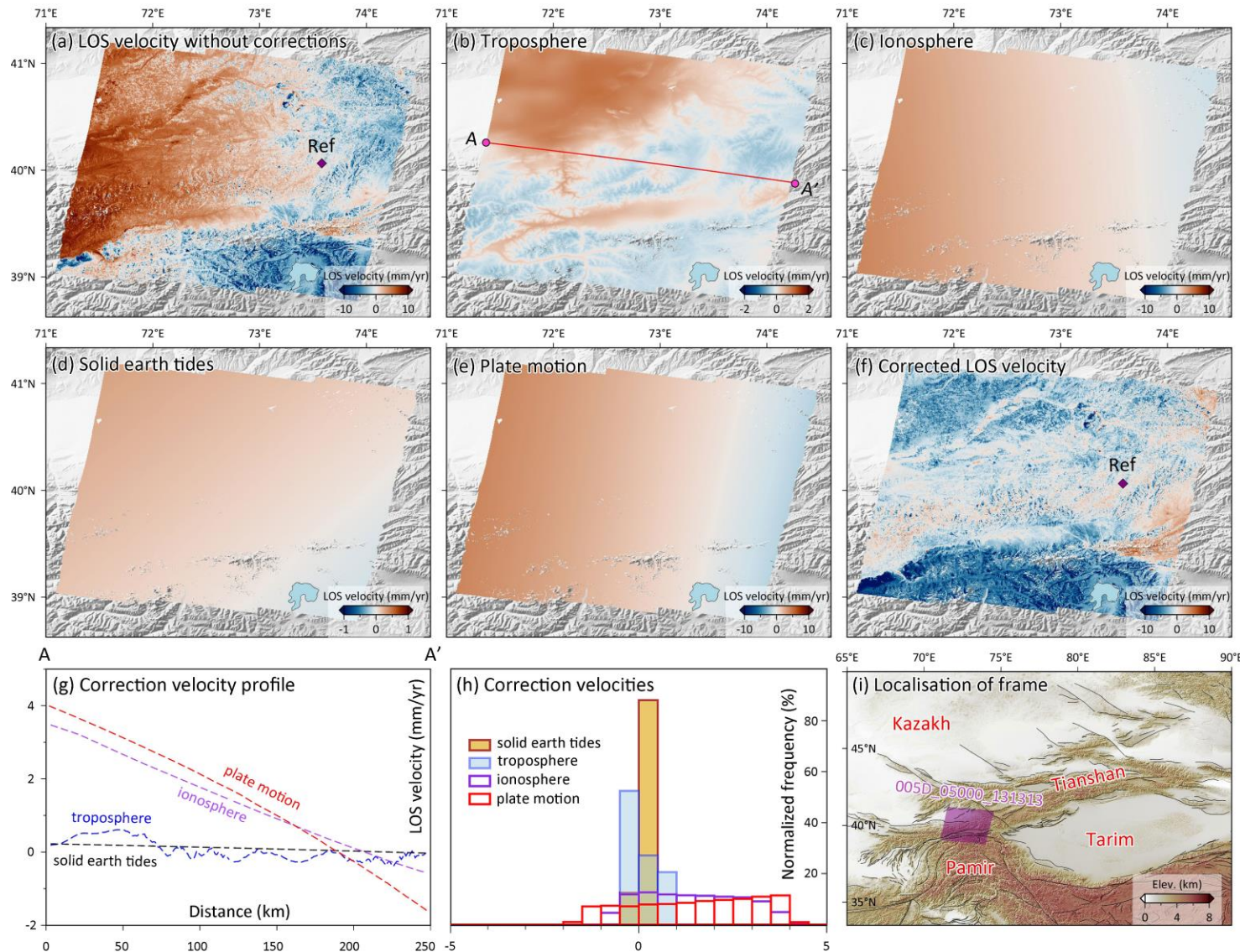
Time series



(Elliott et al., 2026)

Real-time standard interferogram products

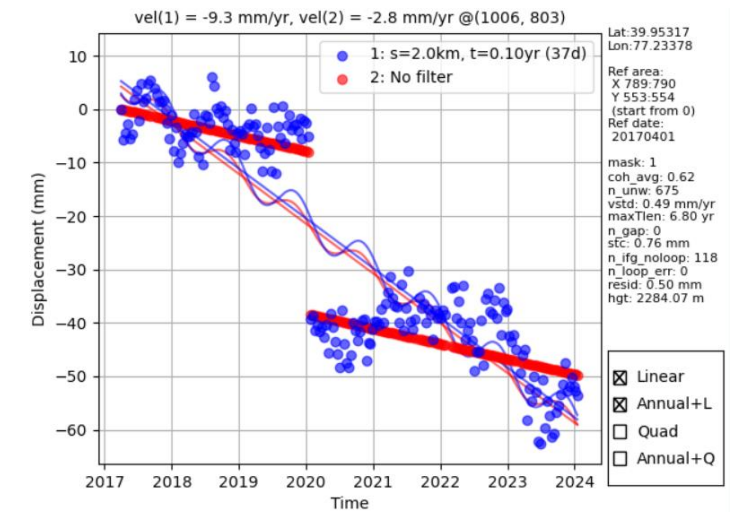
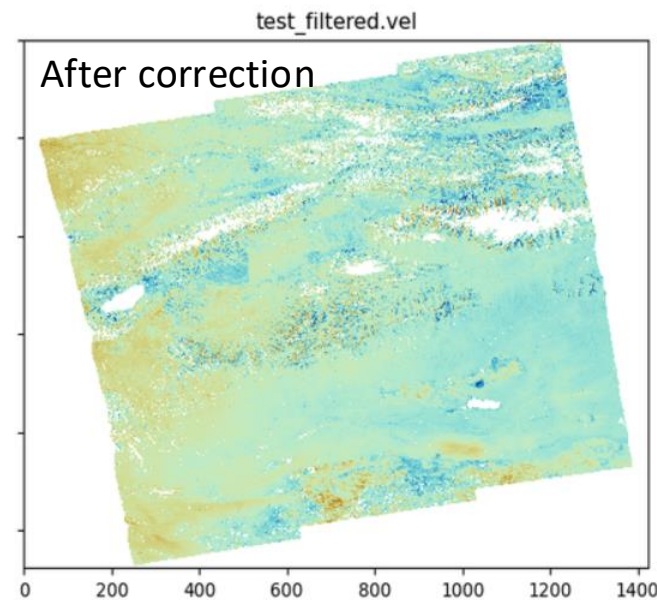
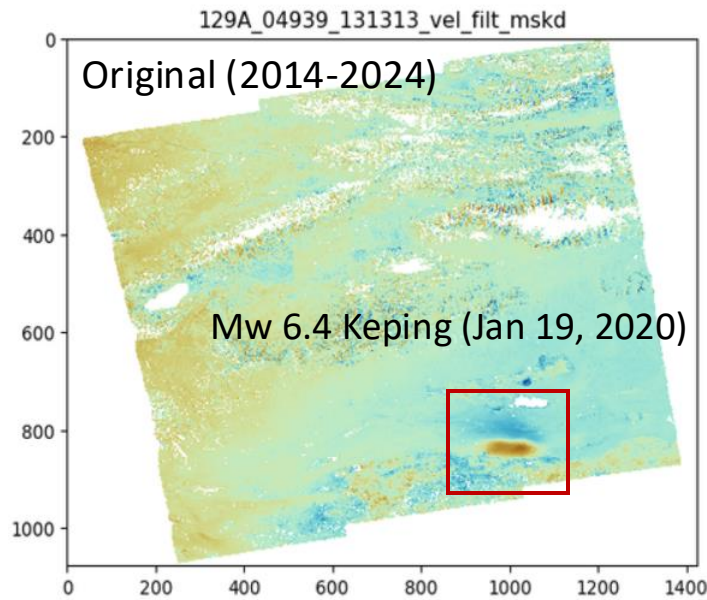
Correction of Non-Tectonic Errors



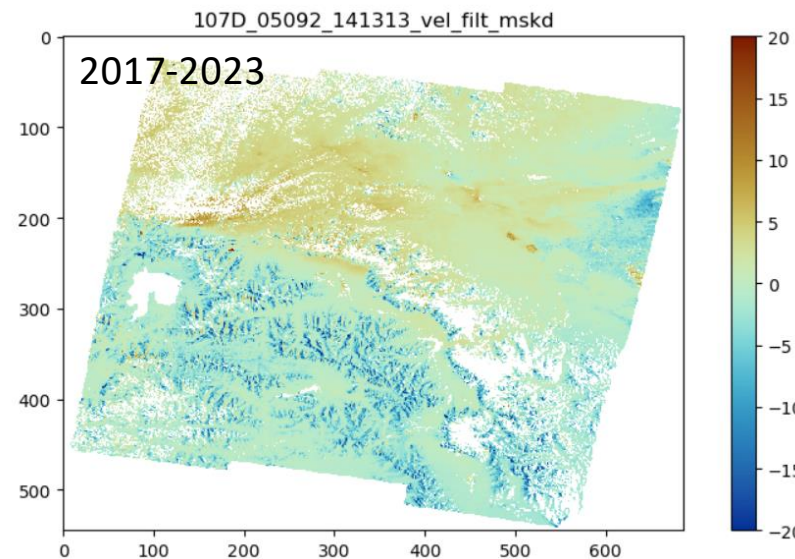
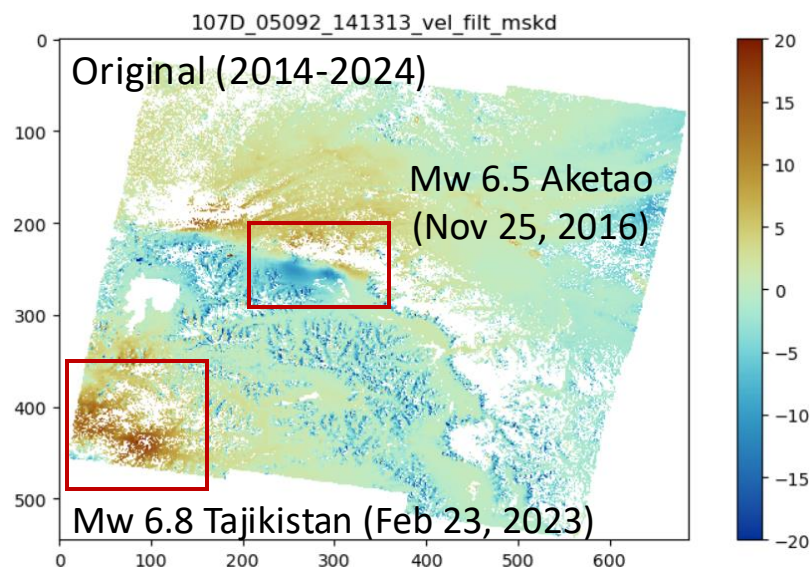
- Troposphere (GACOS) (Yu et al., 2018)
- Ionosphere (Zhang et al., 2022)
- Solid earth tide (Earth tide tool in GMT) (Wessel et al., 2019)
- Plate motion (Stephenson et al., 2022)

Application of corrections on example frame 005D_05000_131313

Co-seismic deformation correction



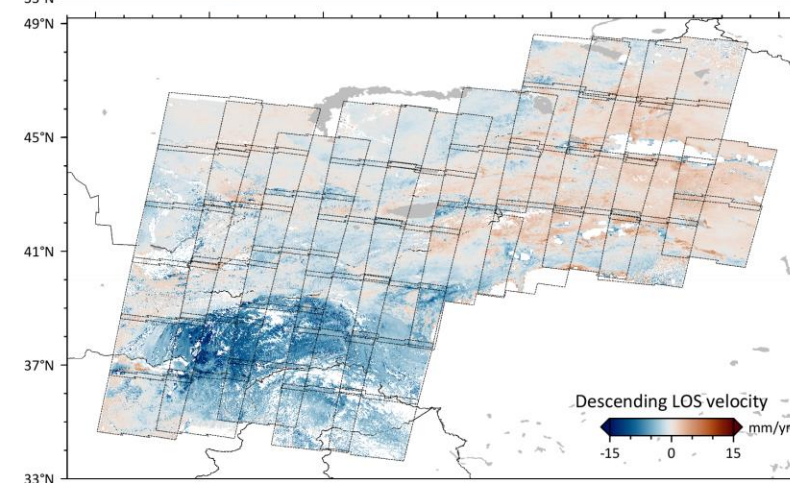
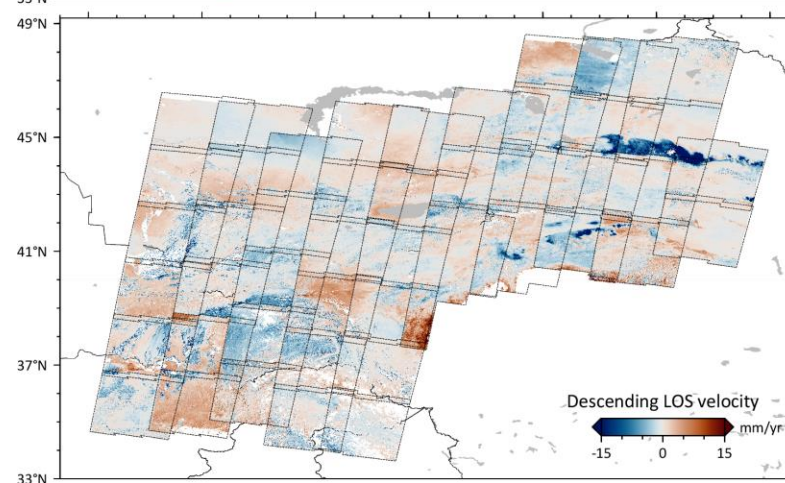
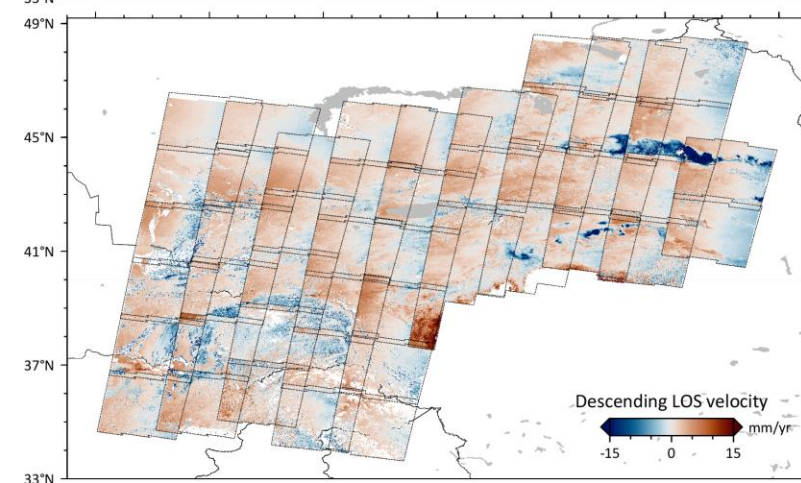
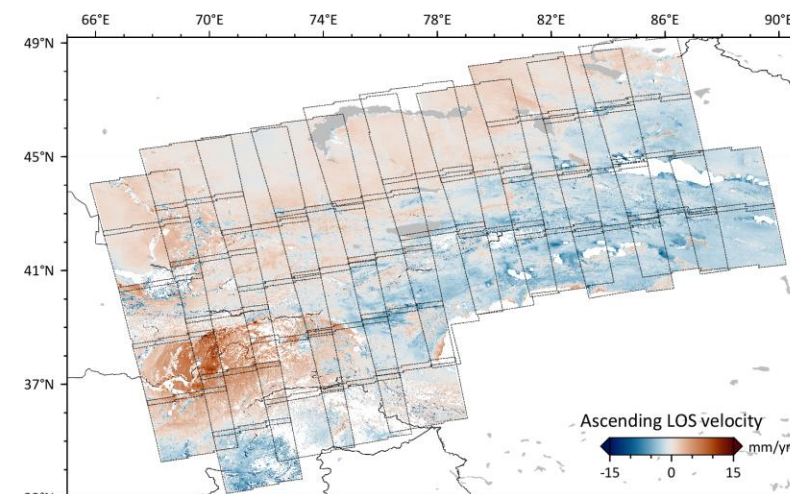
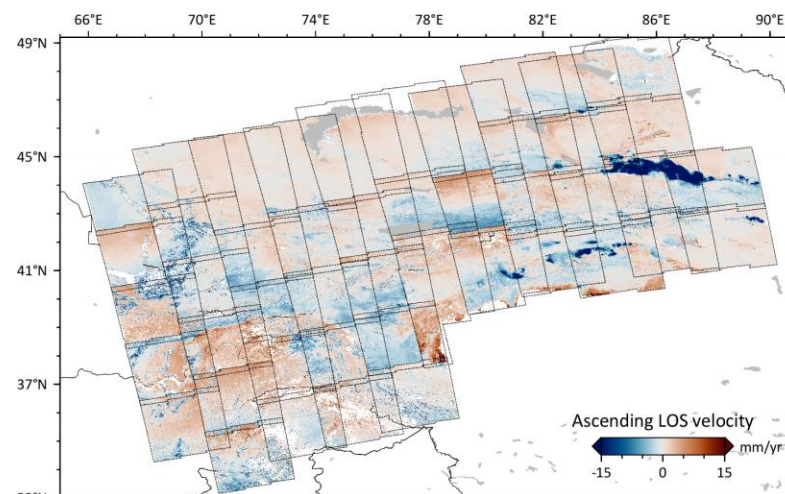
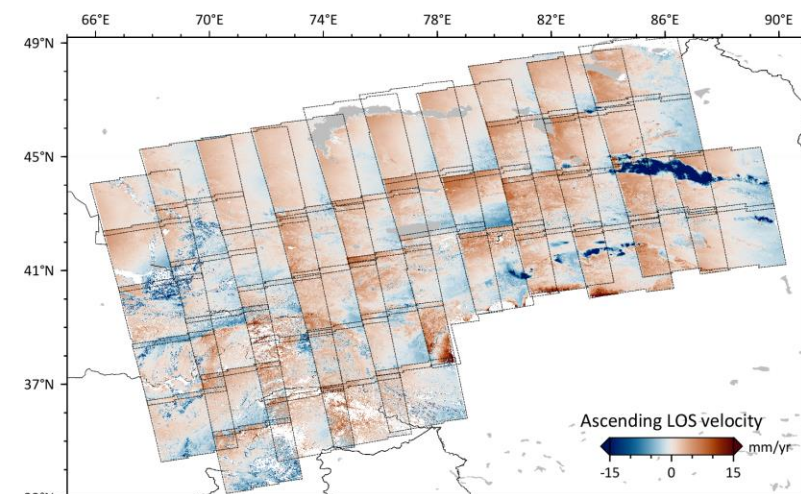
Fitting a Step Function (pixelwise)



Time Series Truncation
(After/Before Earthquake)

6 earthquakes, 21 frames

Line-of-Sight Velocities

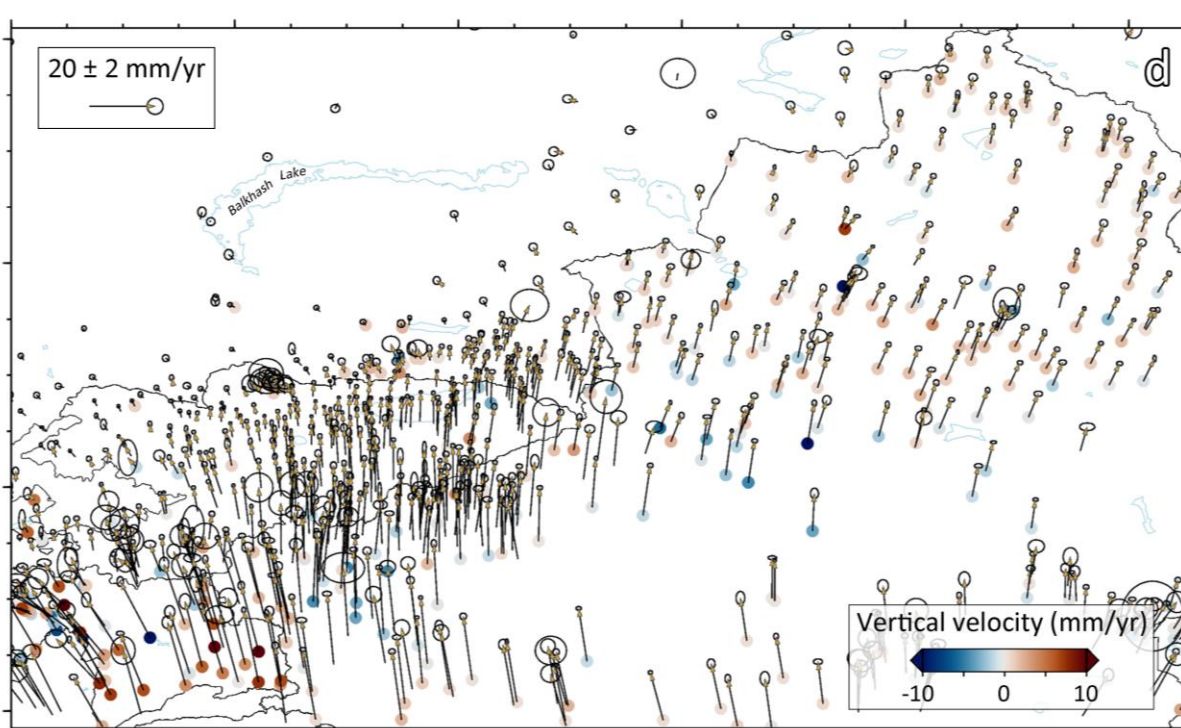


LOS Velocities
(Referenced separately)

LOS Velocities after corrections
(Referenced separately)

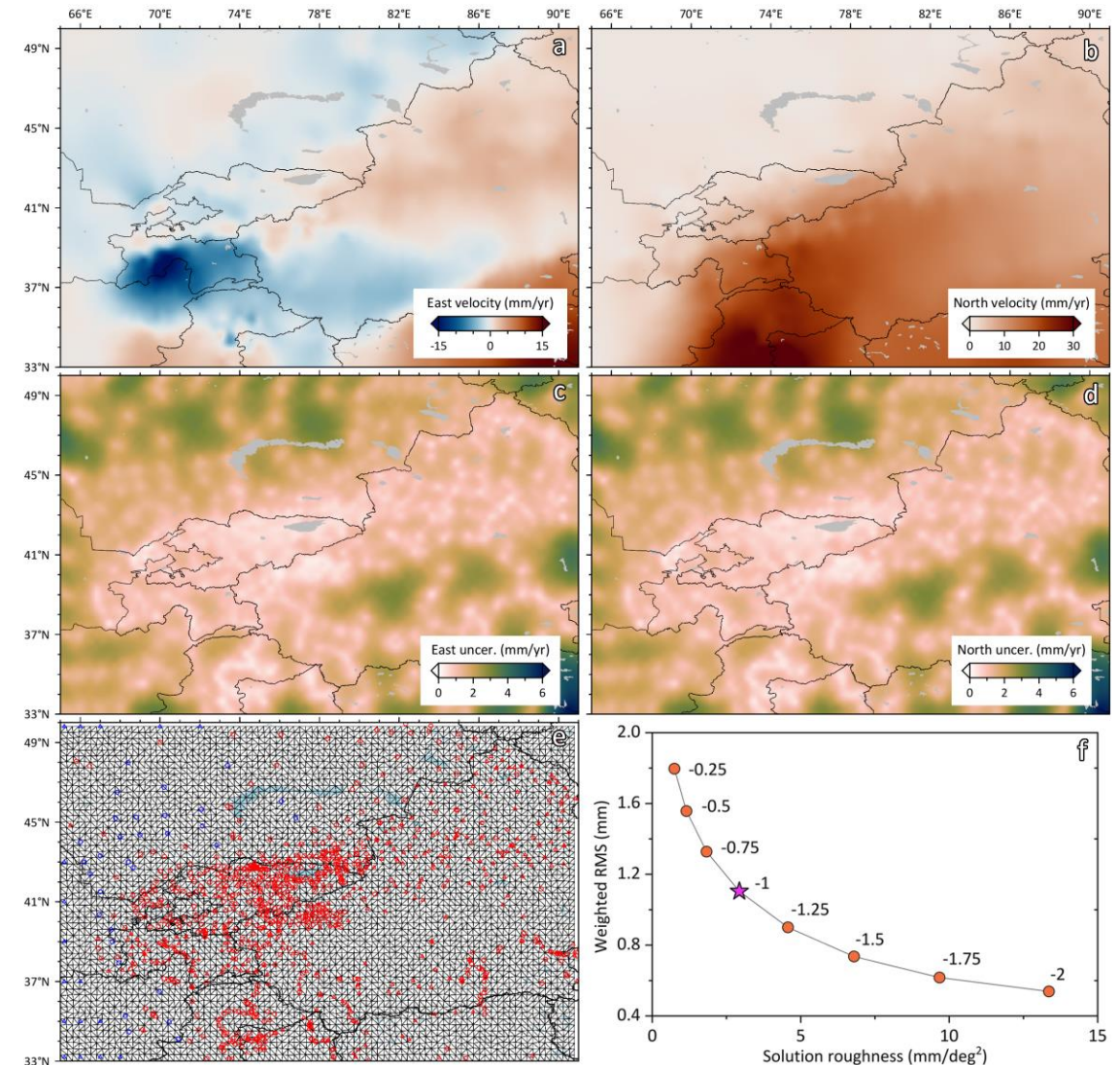
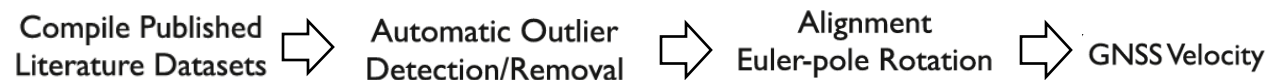
Eura-fixed LOS
(Align with GNSS reference frame)
Mask threshold (15 mm/yr)

GNSS Velocity Fields



~1,000 GNSS velocity vector (2D+3D) with respect to the ITRF2014 Eurasia-fixed reference frame.

From AHB GNSS Dataset ([Elliott et al., 2026](#))

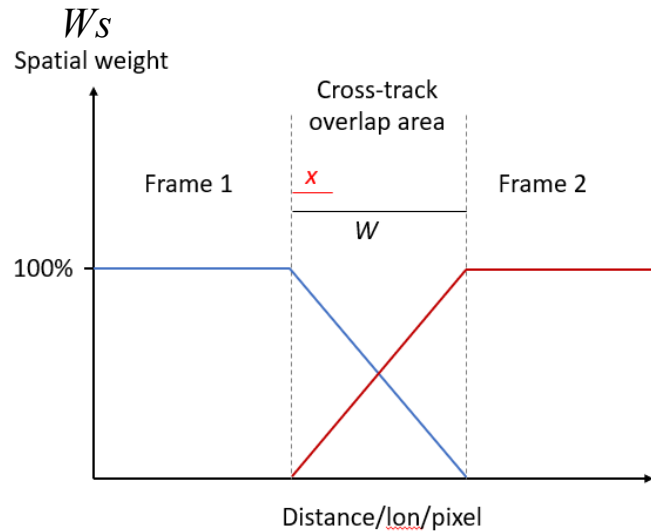


VELMAP Interpolation ([Wang & Wright, 2012](#))

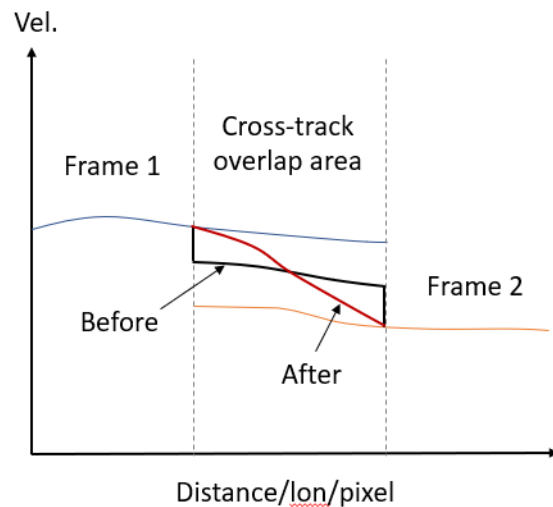
Velocity Decomposition



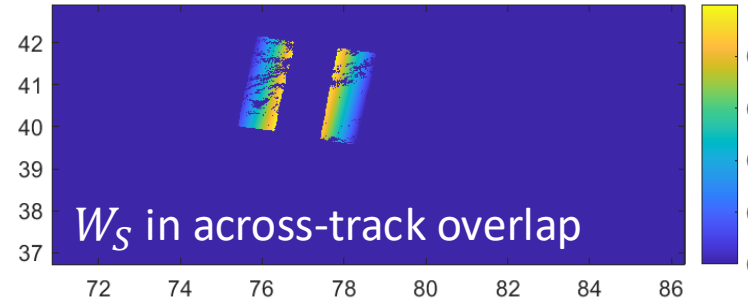
Asc. + Des. + GNSS north $\xrightarrow{\text{WLS}}$ East and Vertical



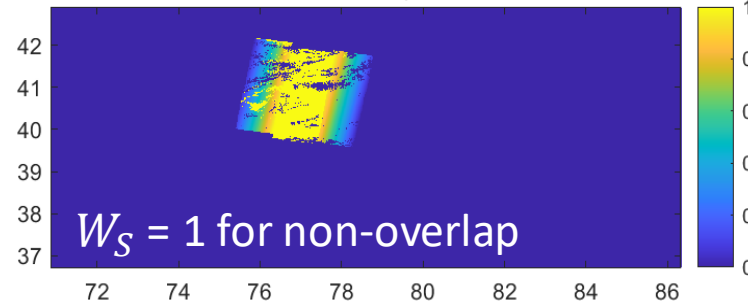
$$\omega_L(x) = \frac{W - x}{W}, \quad \omega_R(x) = \frac{x}{W}$$



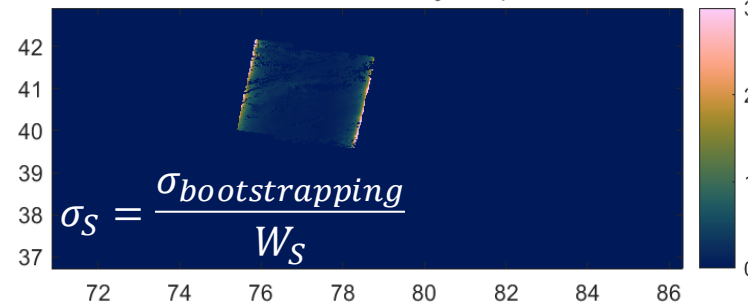
Overlap-only spatial weight for 034D₀4913₁31313 (descending)



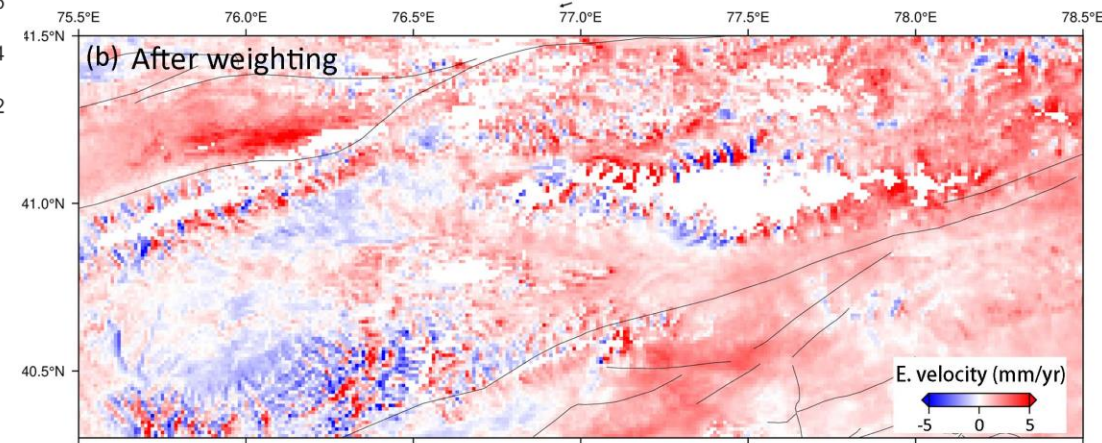
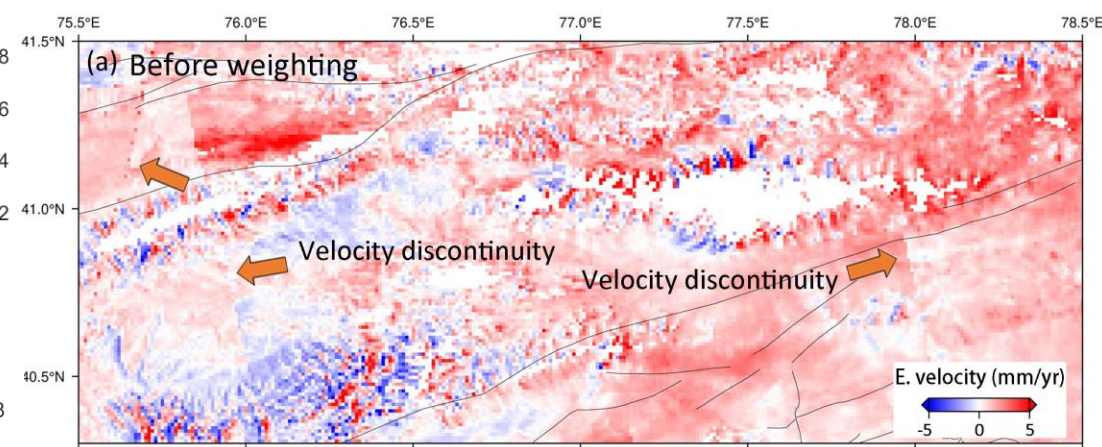
Full-frame spatial weight for 034D₀4913₁31313 (descending)



vstd after weight for 034D₀4913₁31313

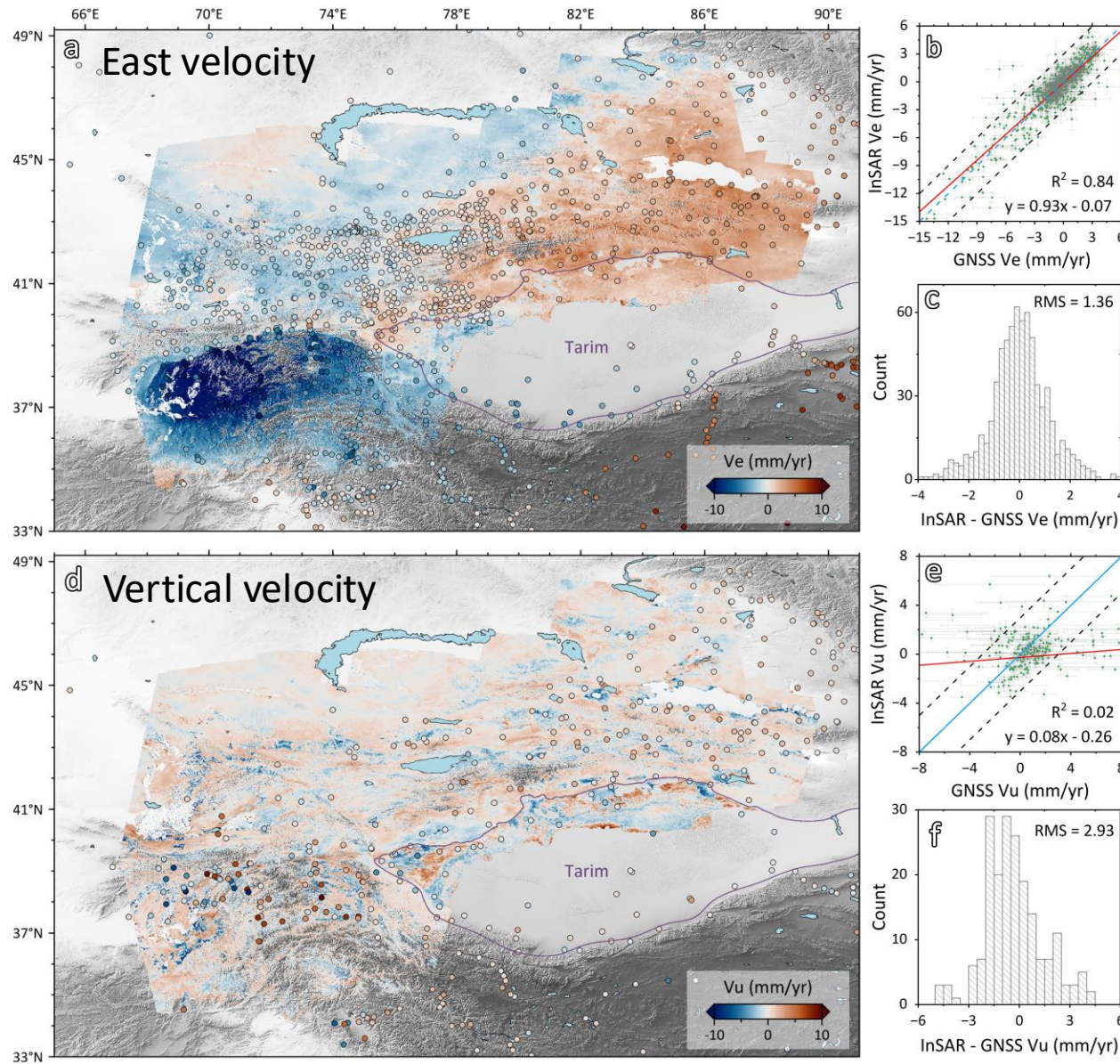


Combining formal uncertainty ($\sigma_{bootstrapping}$) with W_S



Comparison of velocity before and after spatial weighting

Interseismic Velocity



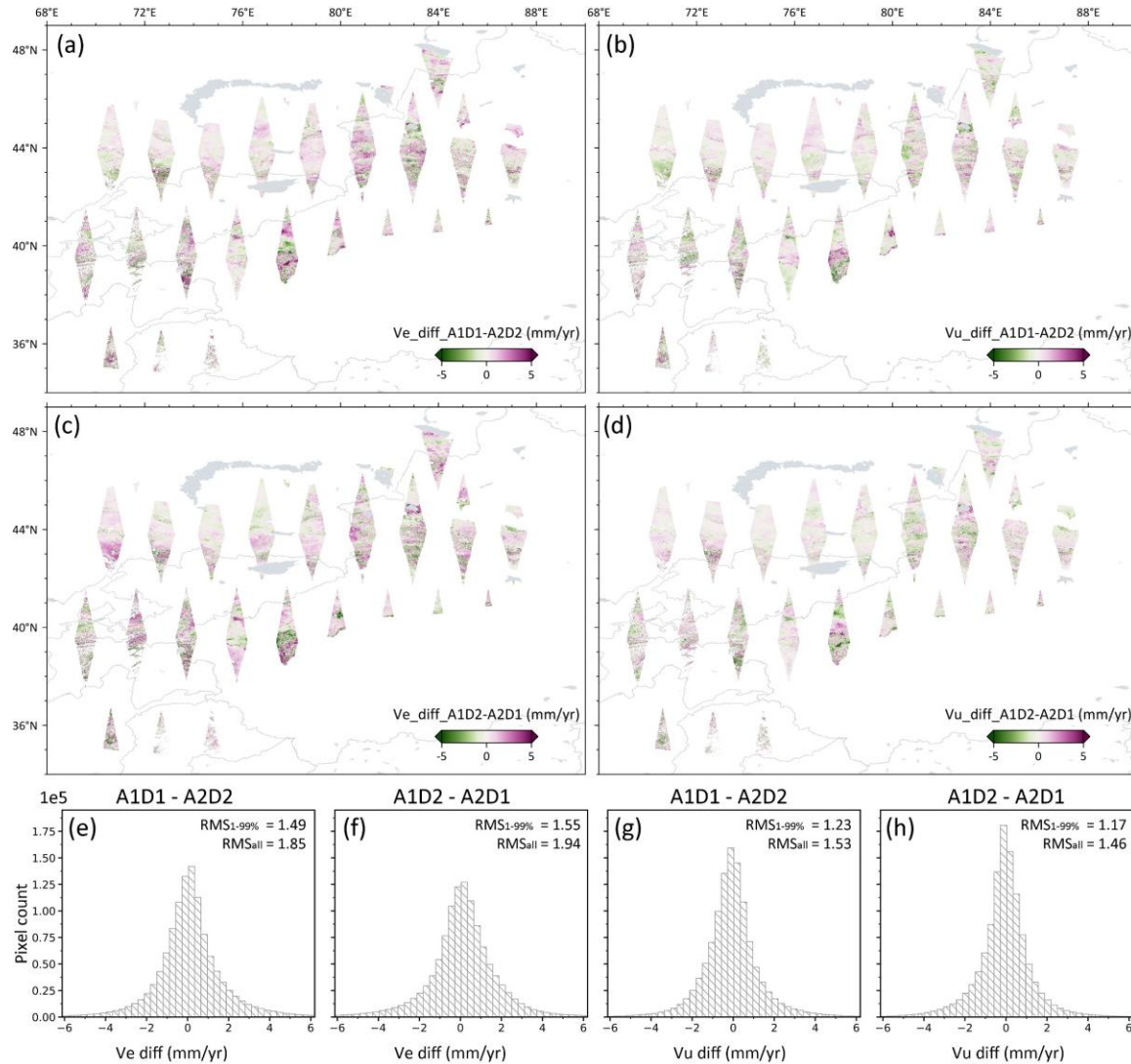
500 m resolution, ~2 million km²

- East-west velocity shows good agreement with GNSS
- Vertical velocity is more difficult to interpret
- Vertical velocity is suppressed when aligned with GNSS

Uncertainty Estimation from Redundant Observations



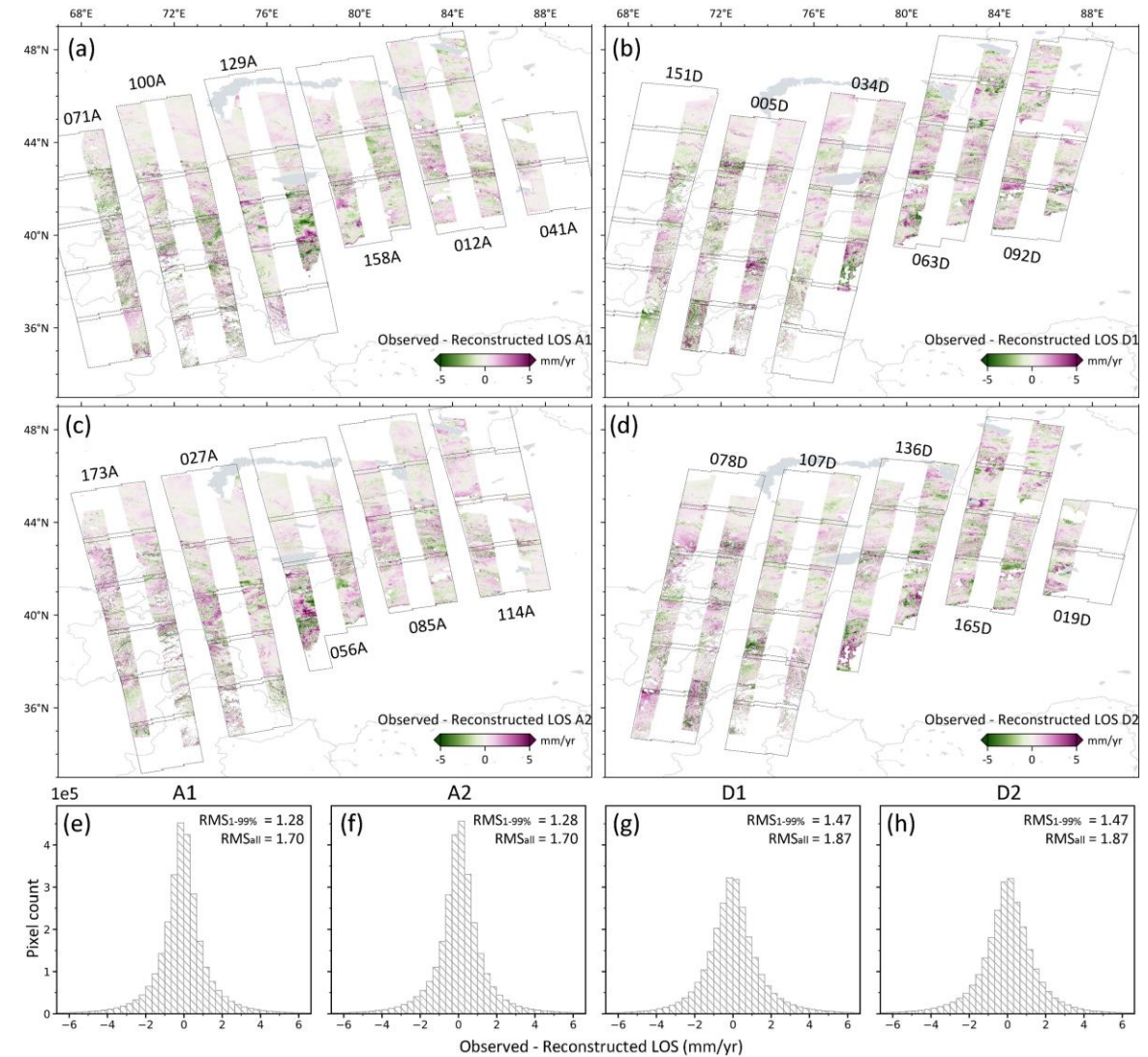
4 independent LOS observations



sig $V_e \approx 1.1$ mm/yr

sig $V_u \approx 0.9$ mm/yr

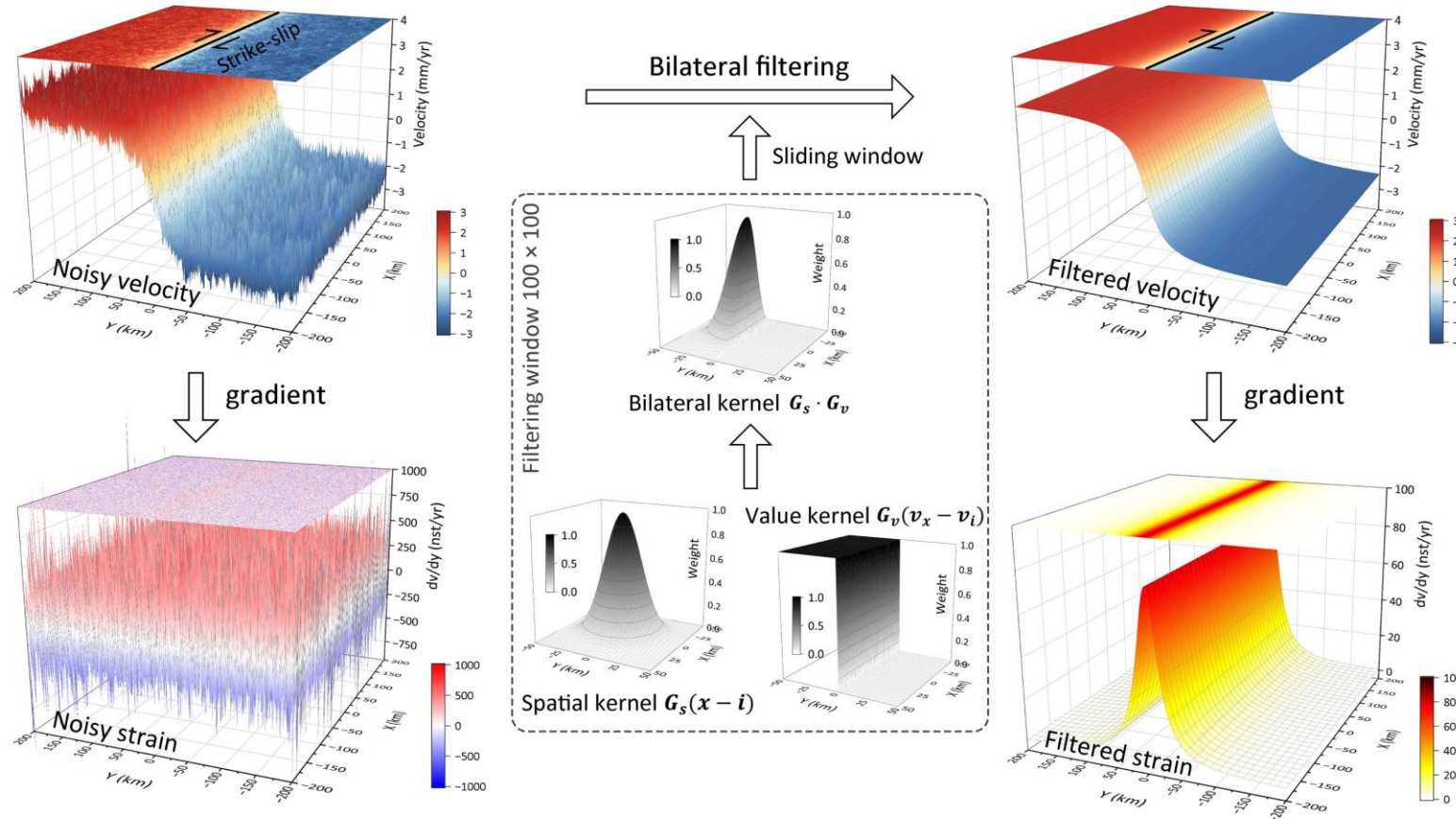
3 independent LOS observations (Wright et al., 2026; Elliott et al., 2026)



sig $V_{los_a} \approx 0.8$ mm/yr

sig $V_{los_d} \approx 1.1$ mm/yr

Strain Rate Calculation – Velocity filtering



Development of **new InSAR velocity-field filter** that better recovers both magnitude and localization of on-fault strain rate.

$$G_s = \exp\left(-\frac{\|x - i\|^2}{2\sigma_s^2}\right)$$

Spatial kernel

$$G_v = \exp\left(-\frac{\|v(x) - v(i)\|^2}{2\sigma_v^2}\right)$$

Value kernel

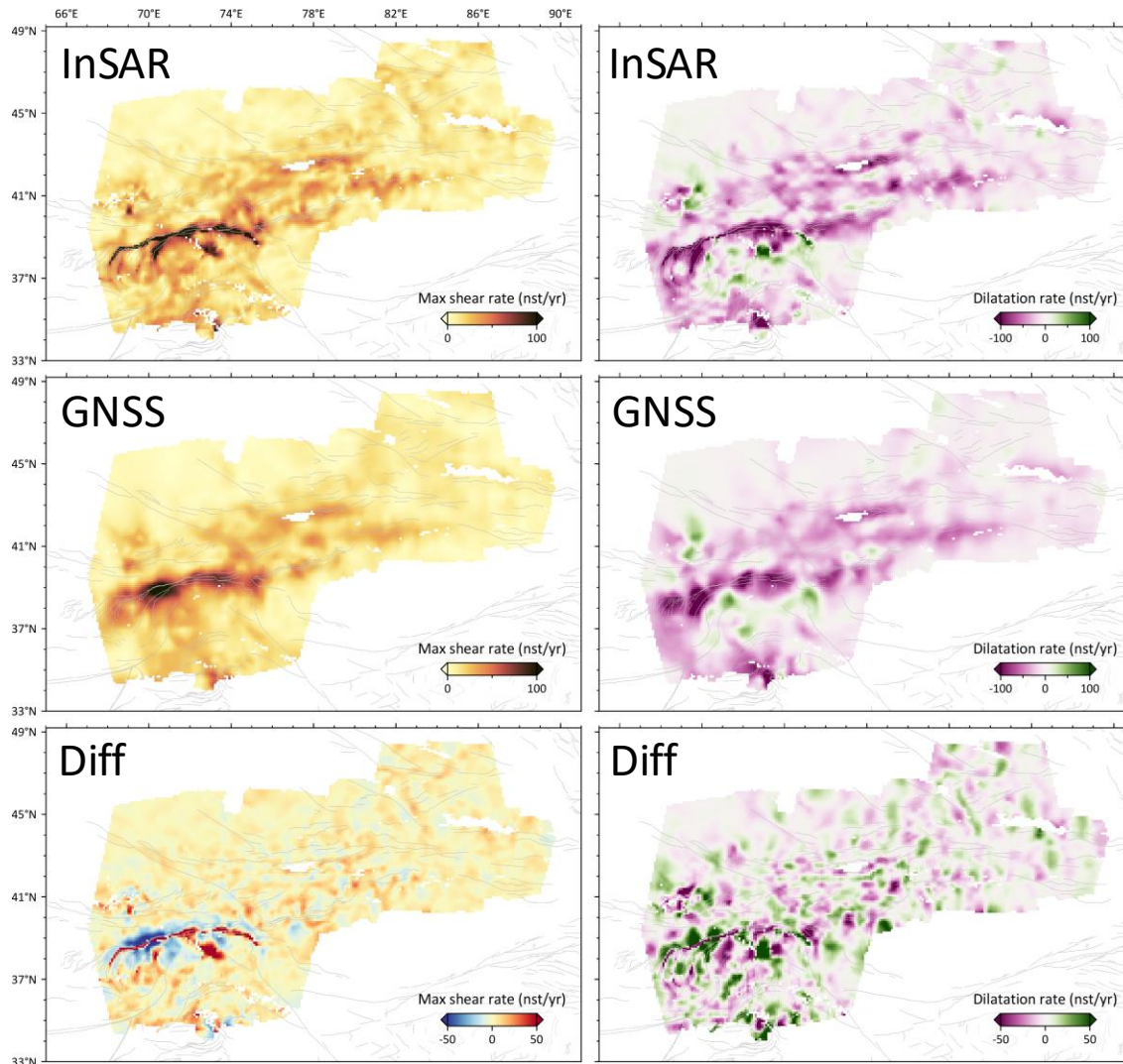


Schematic diagram of the bilateral filtering method and its influence on strain rate estimation

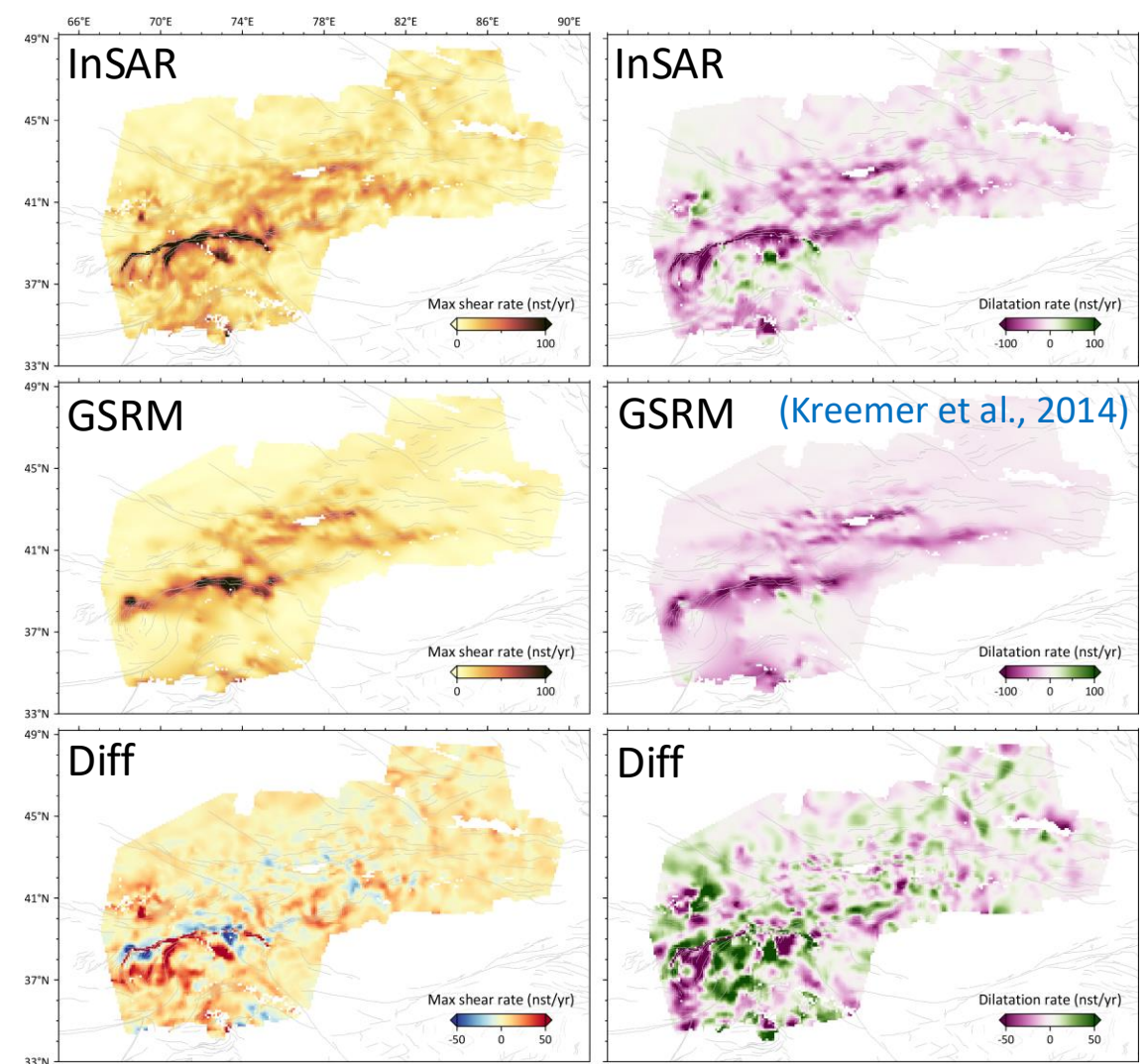
POSTER SESSION II (Chang et al., 2026)

$$v_{\text{filtered}}(x) = \frac{1}{W_p} \sum_{i \in \mathcal{N}(x)} v(i) \cdot G_s(\|x - i\|) \cdot G_v(\|v(x) - v(i)\|)$$

Strain Rate Comparison



InSAR strain rate compare with **GNSS-only**

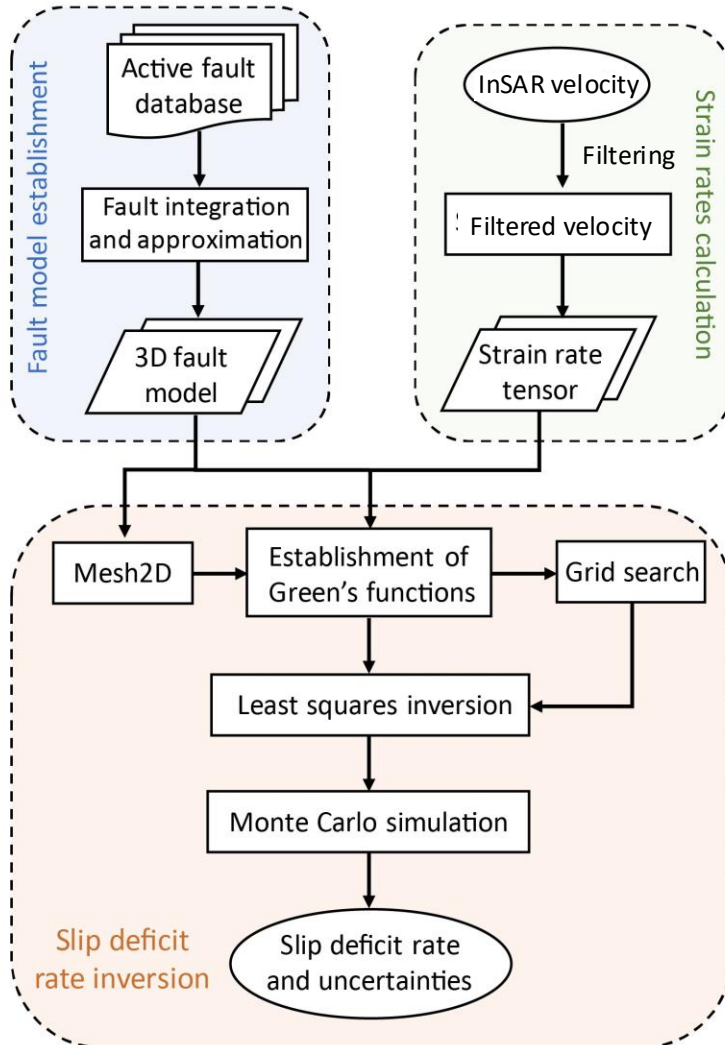


InSAR strain rate compare with **GSRM v2.1**

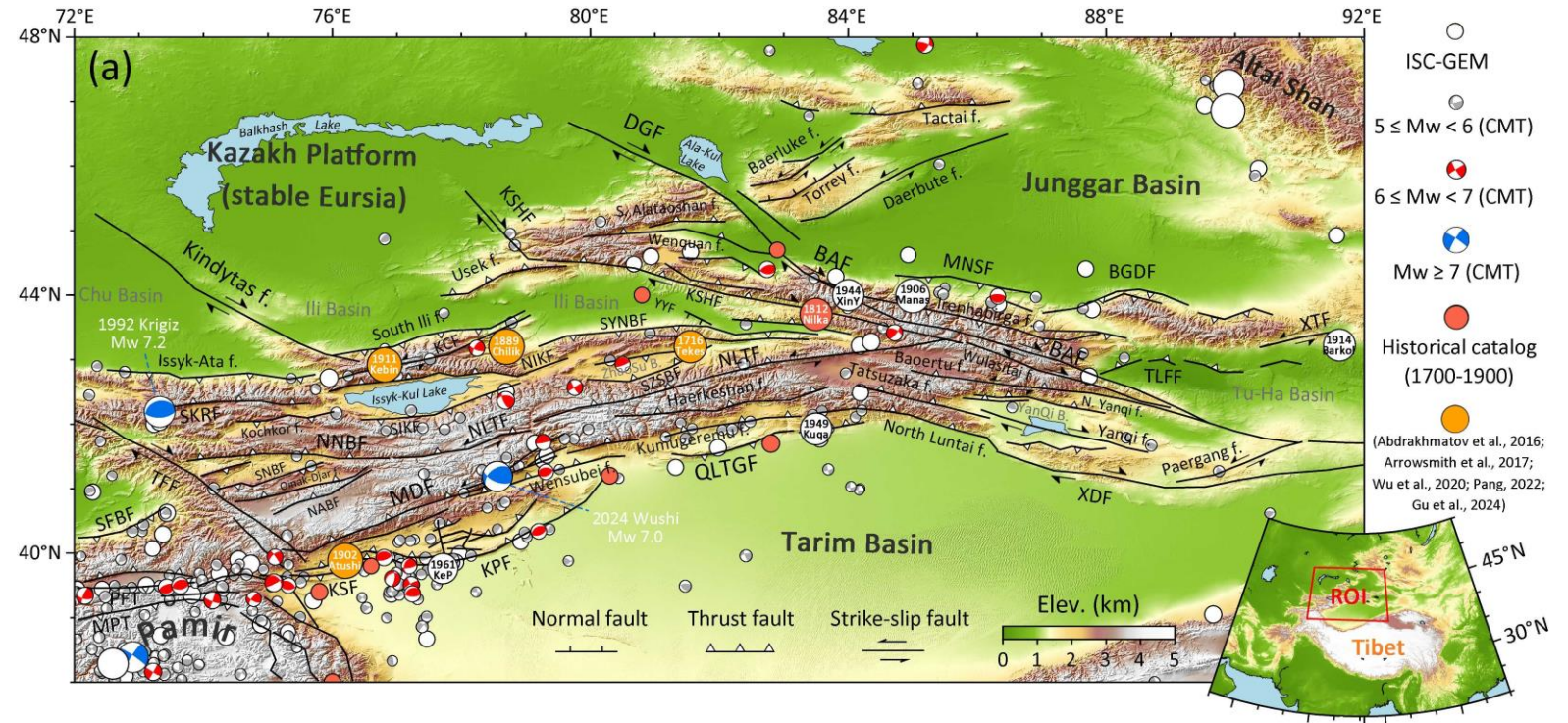
Fault Slip Deficit Rate Inversion



Strain-Rate-Based Inversion Avoids block partitioning, allows for systematic inversion of closely spaced faults (Johnson et al., 2024)



Workflow



(Chang et al., 2025)

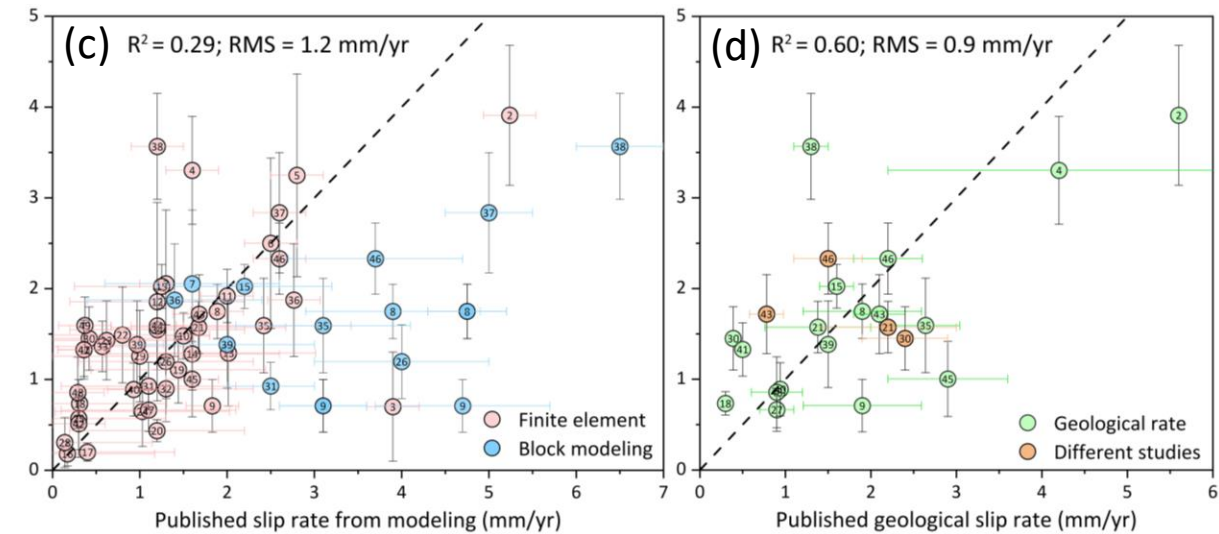
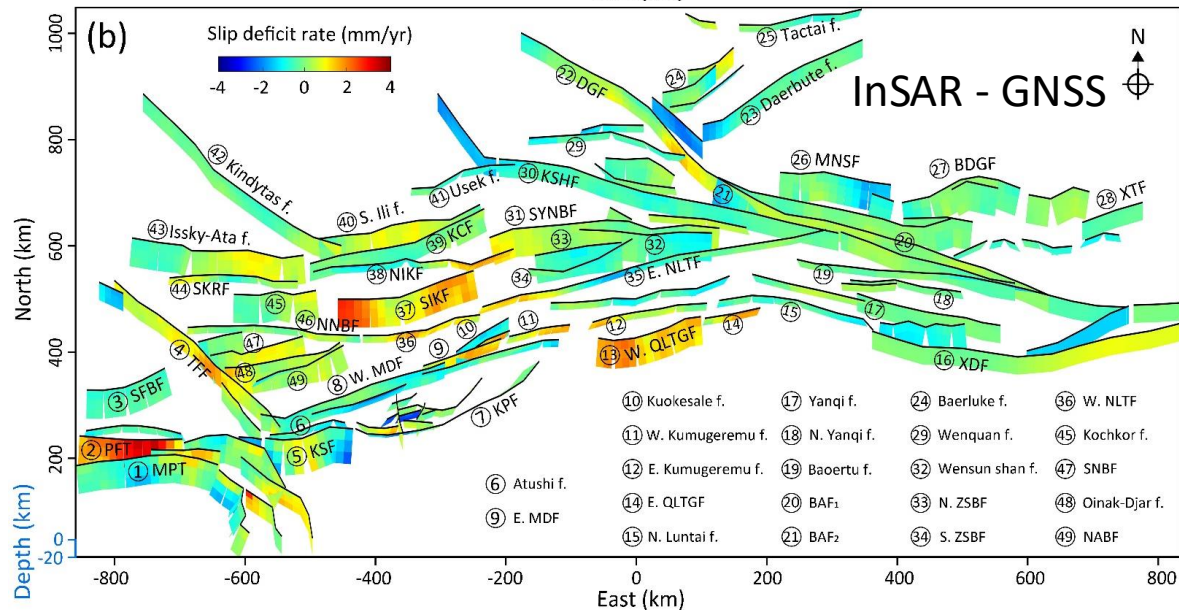
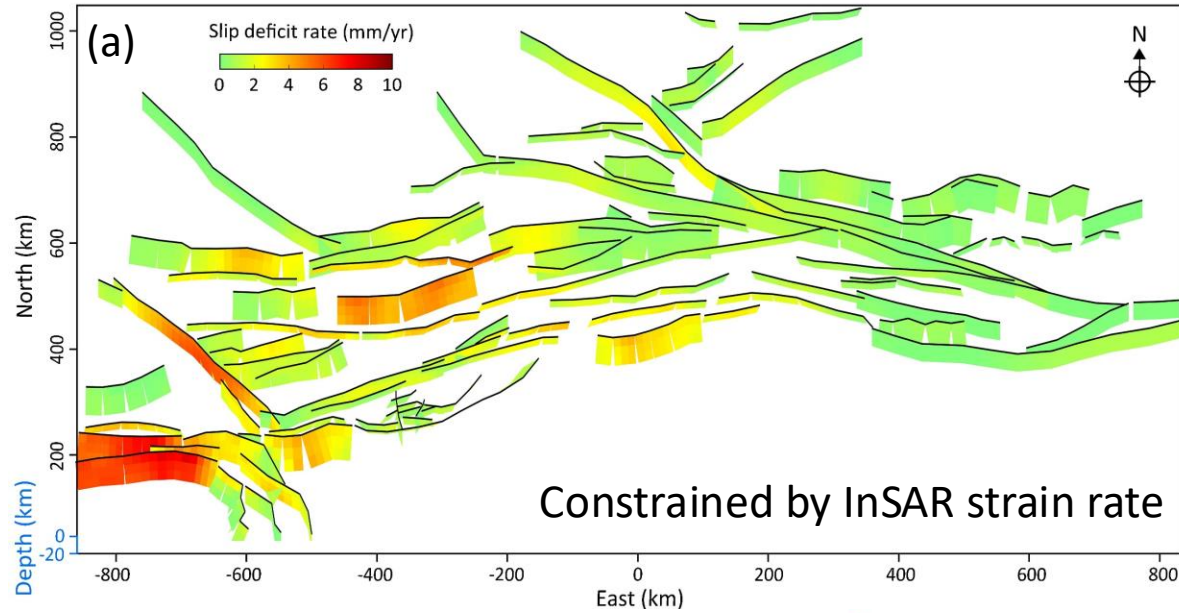
Fault database:

the China Active Faults Database, GEM global active faults database, and published studies.

Earthquake catalog:

ISC-GEM (1904-2024), GCMT (1976-2024), the China Historical Earthquake Catalog (1700-1900), and paleoseismological studies.

Fault Slip Deficit Rate



- InSAR strain rates reveal more details about fault slip deficit.
- Discrepancies may arise from differences in fault system simplifications, spatial limitations in geological rate measurements, timescale dependent variations in fault behavior.
- ~60% of strain rate is accommodated by major faults (interseismic coupling), while the remaining is associated with distributed moment sources (diffuse crustal deformation).

Seismic Moment Accumulation



Assumptions:

Elastic energy accumulated since the last major earthquake.

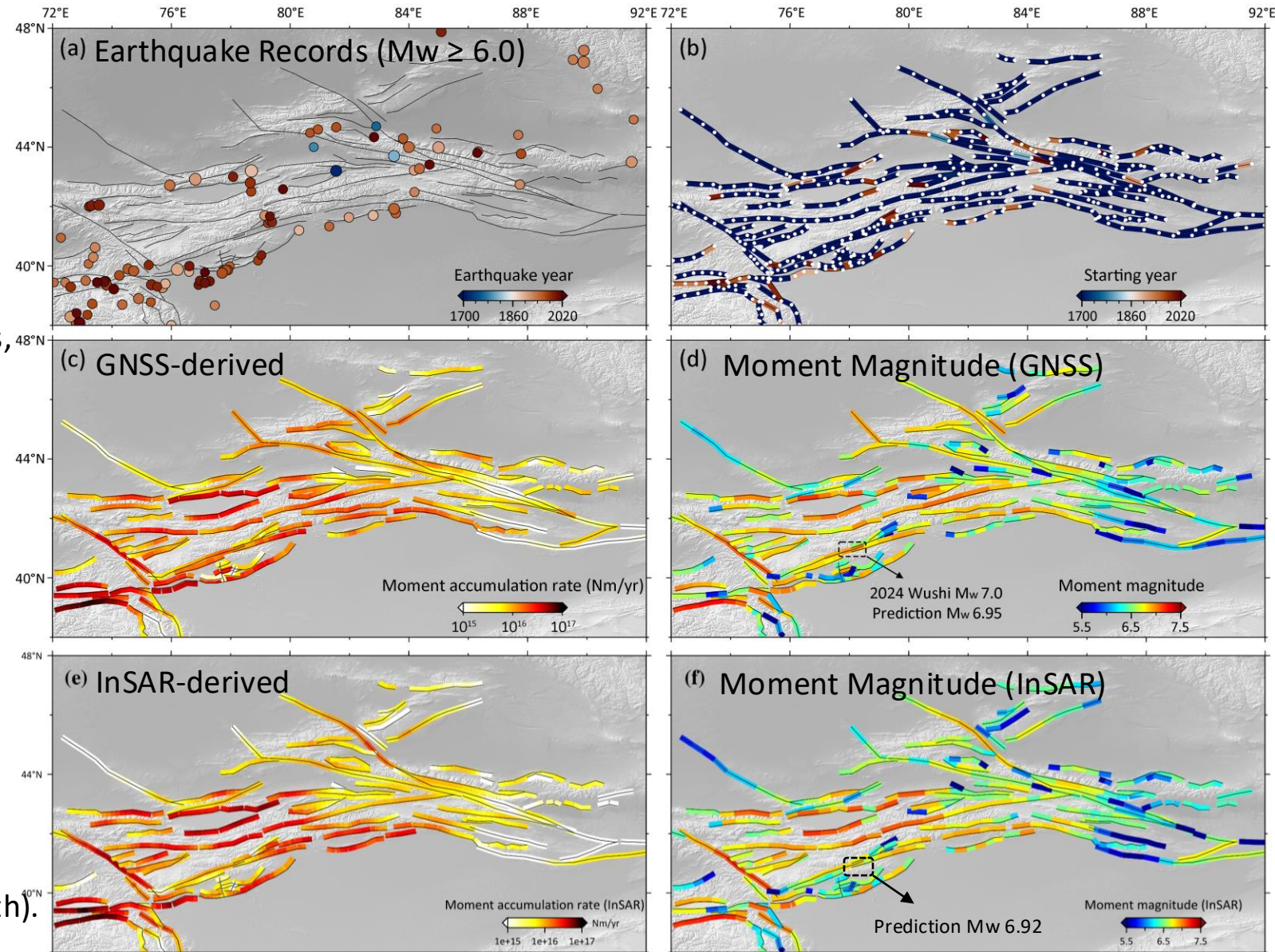
For segments without documented seismic events, we set the accumulation start time to 1700 (lower bound).

For earthquakes lacking precise location constraints, we associate them with the nearest mapped fault.

20 out of 439 fault segments are capable of generating $M_w \geq 7.0$ earthquakes.

Uncertainties:

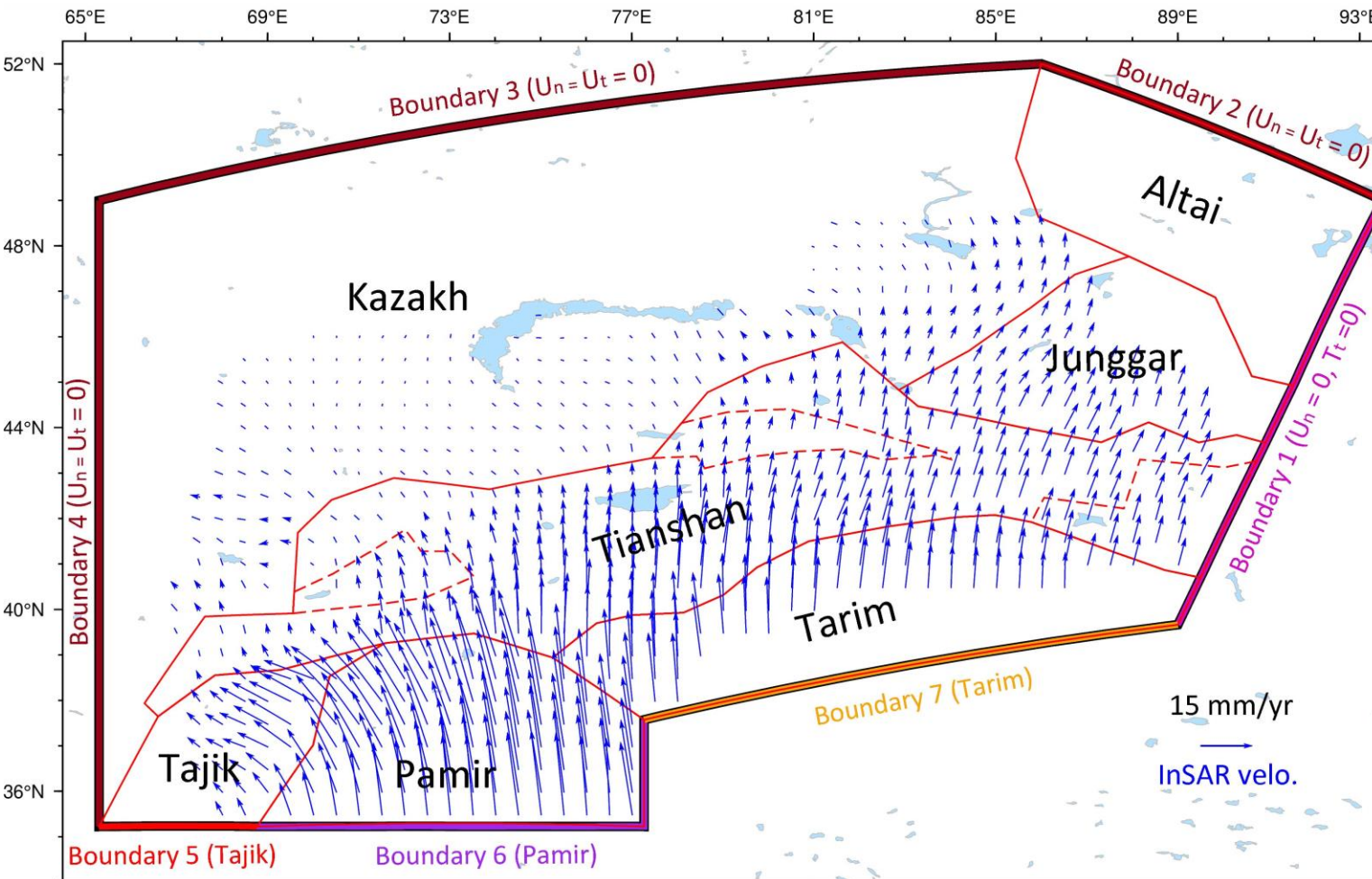
- Simple fault geometry (segmentation, dip, rake).
- Incompleteness of historical earthquake records.
- Fault intersection, cascading ruptures.
- Strain uncertainties from GNSS Interpolation (north).



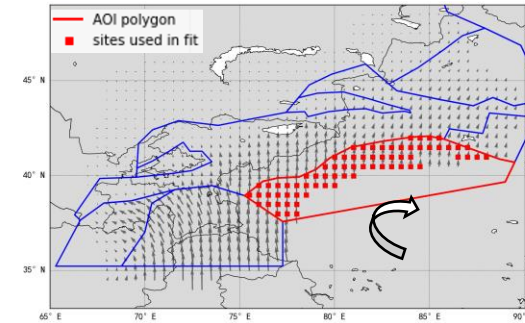
What controls the present-day deformation of Tianshan? 2D (Vertical averaged) Viscous Shell Model

(Houseman et al., 1986)

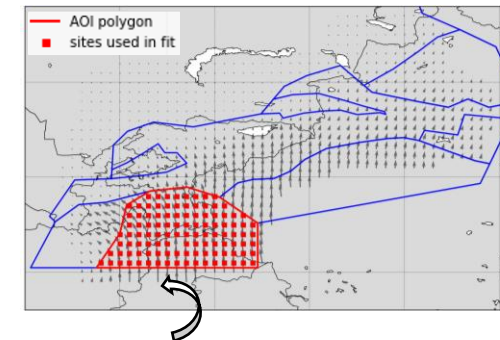
Downsample to 50 km (836 observations)



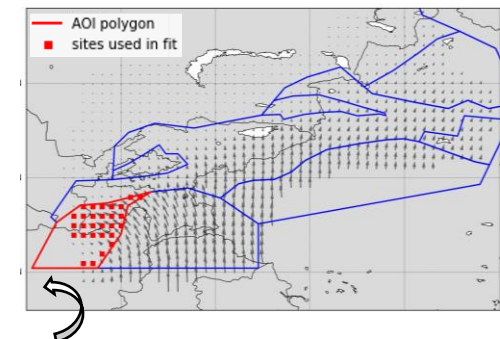
Boundary Conditions



Tarim boundary:
Euler pole: 95.7E 38.7N
-0.68 deg/Myr



Pamir boundary :
Euler pole: 59.5E 33.0N
0.84 deg/Myr

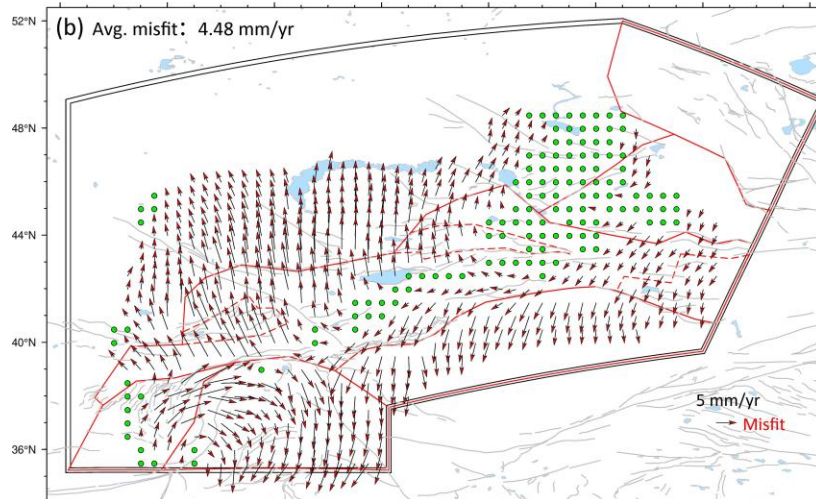
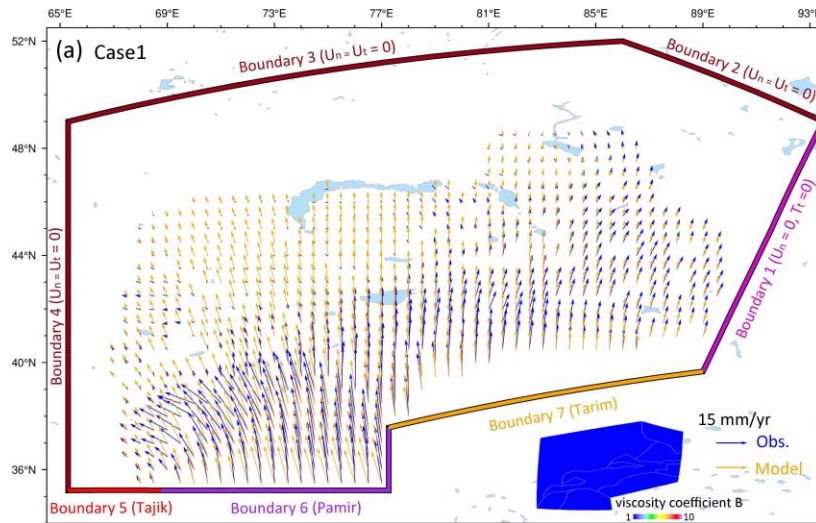


Tajik boundary:
Euler pole: 66.4E 34.8N
1.65 deg/Myr

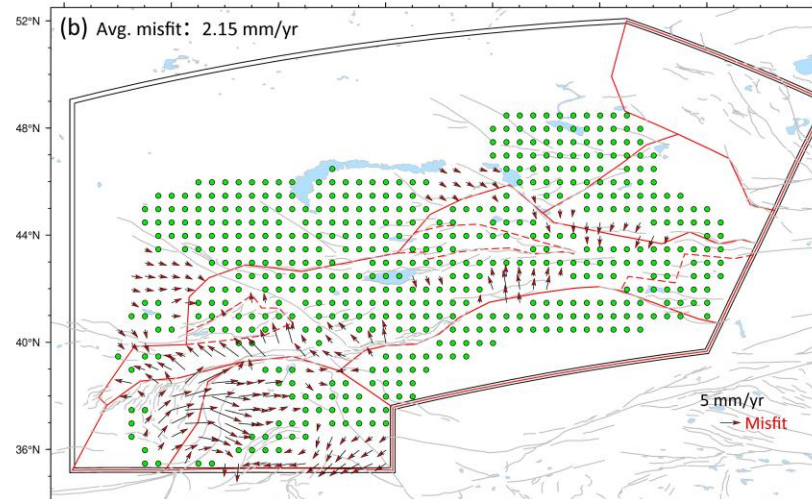
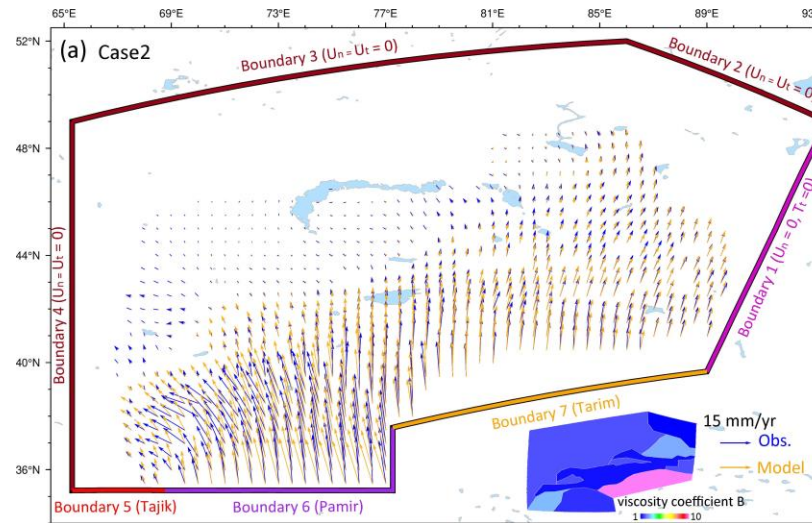
Different Model Settings



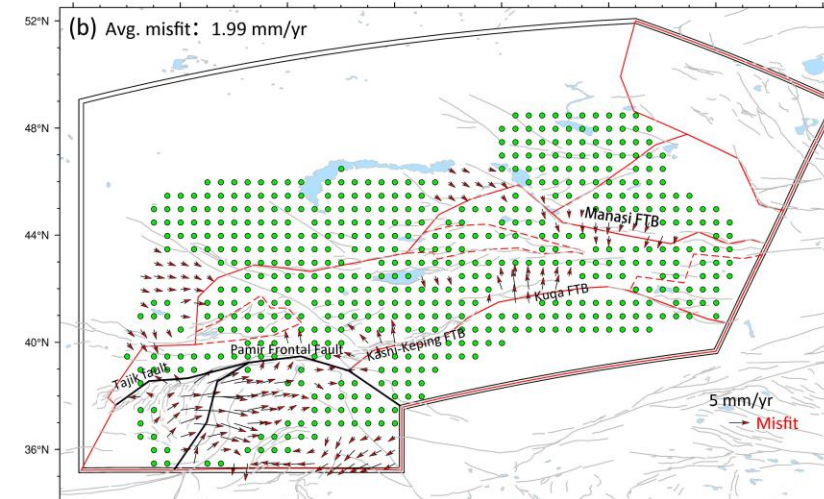
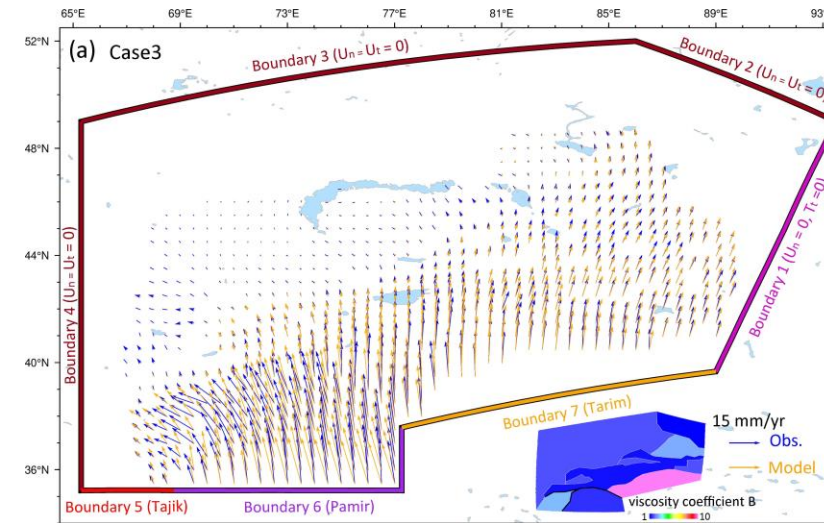
Case1: Uniform strength



Case2: Variable strength (Optimal Param.)



Case3: Add fault (ft and fn)

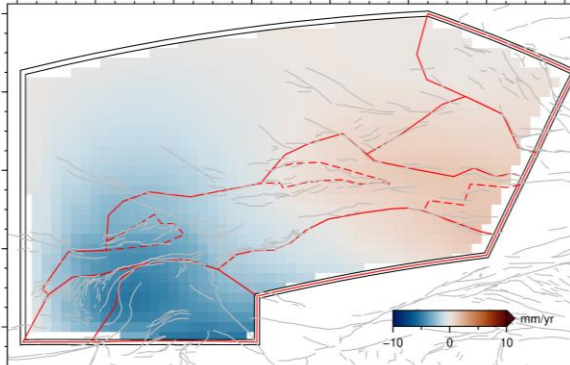


Model Comparison

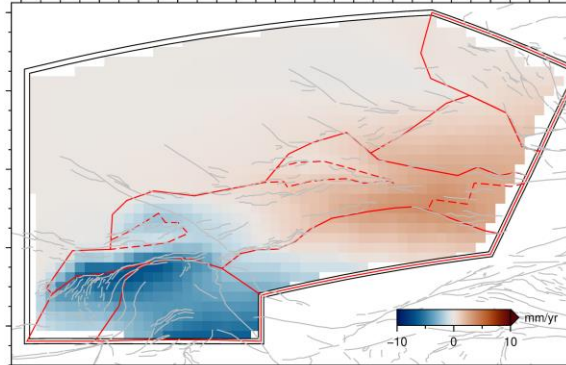


East velocity

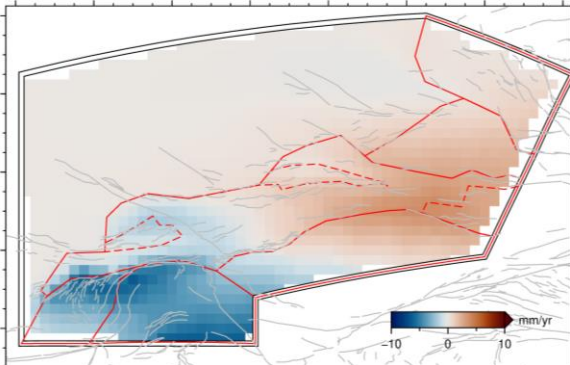
Case1 (uniform strength)



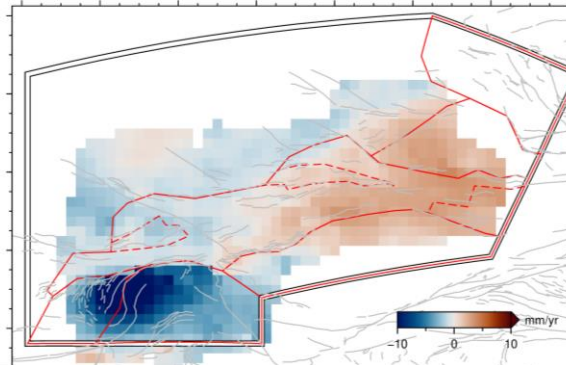
Case2 (variable strength)



Case3 (Add fault)

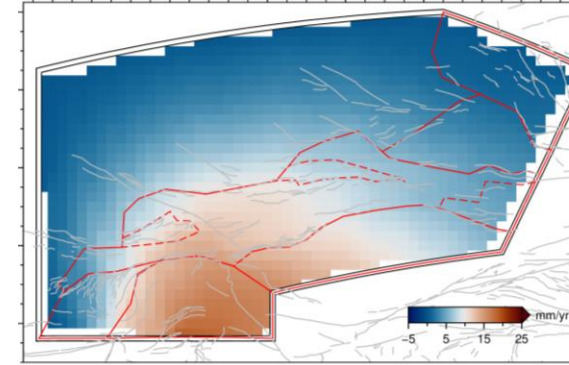


Observation

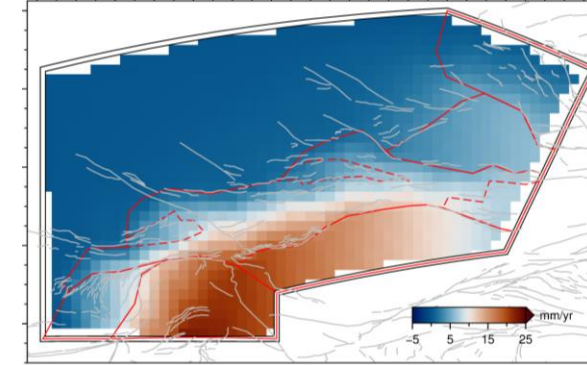


North velocity

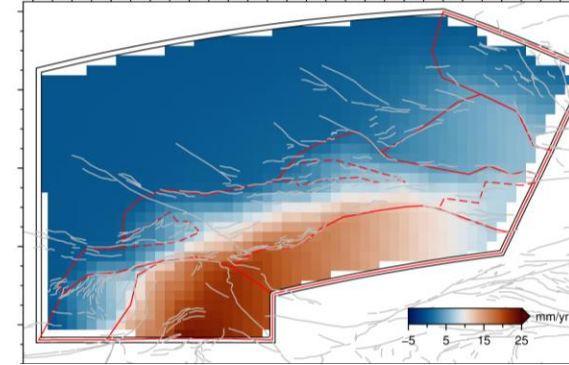
Case1 (uniform strength)



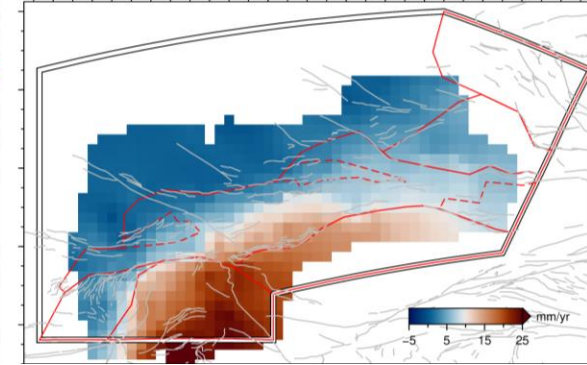
Case2 (variable strength)



Case3 (Add fault)

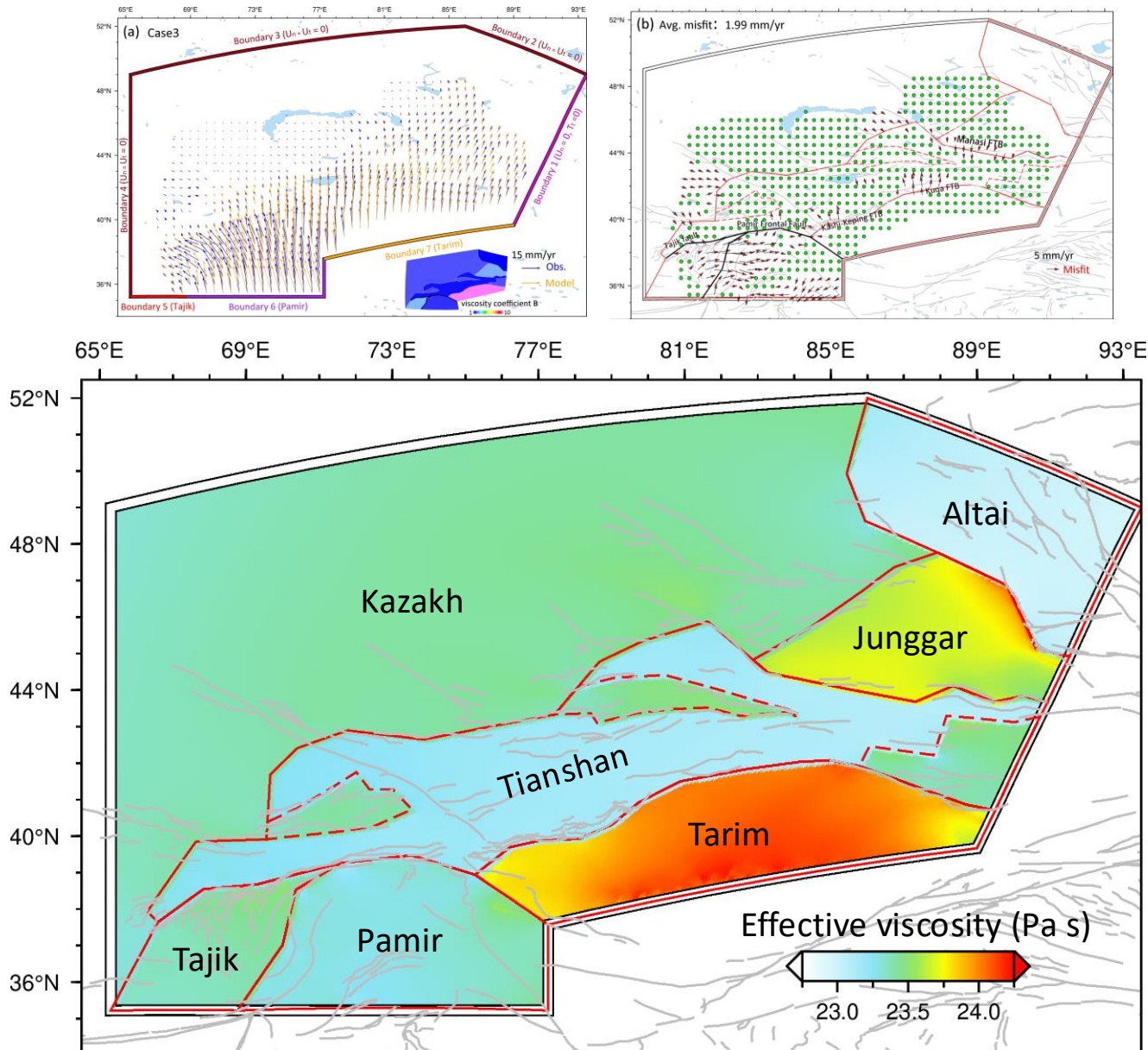


Observation



Main misfits are concentrated in the Pamir and Tajik

Deformation Dynamics



- Indentation of the Tarim and Pamir (boundary loading) is the first-order drivers of present-day deformation in the Tianshan.
- A relatively weak Tianshan lithosphere (depth-averaged $10^{23}\text{Pa}\cdot\text{s}$) is required to reproduce the observed deformation pattern; the effective viscosity is 2–10 times lower than that of the surrounding Kazakh and Tarim.

Next step:

- GPE is not a necessary factor for explaining the present-day deformation, further investigation is needed.
- The larger misfit near the Pamir-Tajik may be related to the plate subduction process; consider adding additional horizontal traction to the model.
- Further increase the model complexity and observational resolution to investigate rheological differences within the Tianshan.

Take-home messages



- A high-resolution three-dimensional interseismic surface velocity field covering approximately 2 million km² of the Tianshan and its surrounding regions is constructed, providing a data foundation for a better understanding of continental deformation and seismic hazards.
- Slip-deficit rates of 97 active faults in the Tianshan are systematically inverted from surface strain-rate constraints, and the interseismic elastic moment budget and future seismic potential of these faults are evaluated. 20 active fault segments have accumulated sufficient elastic strain energy to generate earthquakes of $M_w \geq 7.0$.
- Indentation of the Tarim and Pamir is the first-order dynamic drivers of present-day deformation; A relatively weak Tianshan lithosphere is required; gradients in lithospheric GPE are not necessary for explaining the present-day deformation; Weak fault primarily modulates the local velocity gradient.

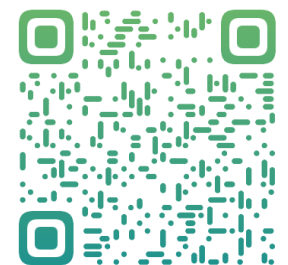


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Actively looking for a postdoc position (Oct. 2026)



(Chang et al., 2025)



(Chang et al., 2026)